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Microstructural analysis of sands with varying degrees of internal stability

J. FONSECA*, W. W. SIM†, T. SHIRE‡ and C. O’SULLIVAN§

Internal erosion involves the migration of particles through a geotechnical structure. Internal erosion poses a significant hazard to embankment dams and flood embankments. The fundamental mechanisms operate at the particle scale and a thorough understanding of these mechanisms can inform the filter design and specification process and reduce the hazard that internal erosion is known to pose to many engineered embankment structures. Engineers have long acknowledged the importance of the grain scale interactions, but until recently, explanations of the mechanisms have been purely hypothetical, as direct observation of the internal structure of filters was not possible. Recent research has used the discrete-element method to establish a particle-scale basis for Kézdi’s filter internal stability criterion. The discrete-element method can provide significant useful data on soil microstructure, so a discrete-element method model is inherently ideal. This study therefore examines a number of real sand samples with varying degrees of internal stability at the particle scale using high-resolution micro-computed tomography. The correlation between coordination number and internal stability is confirmed, with the coordination number values being significantly higher for the real material.

KEYWORDS: erosion; filters; laboratory tests

INTRODUCTION

Embankment dam filters prevent erosion of a low-permeability clay core due to water seepage. Filters should be erosion resistant or ‘internally stable’. Internally unstable filters are susceptible to suffusion, an internal erosion mechanism that occurs when the finer particles within the filter are carried downstream in the direction of seepage flow (e.g. Fell & Fry, 2007). Three factors contribute to internal erosion: (a) material susceptibility, (b) hydraulic susceptibility and (c) mechanical susceptibility (Brown & Bridle, 2009). Design guidelines focus on the material susceptibility and current guidelines for assessing the internal stability of filters are mainly derived empirically and consider the shape of the particle size distribution (PSD) curve (e.g. Kézdi, 1979; Kenney & Lau, 1985, 1986; Burenkova, 1993).

The initiation and propagation of internal erosion is a particle-scale phenomenon. Prior discussions on internal erosion have often included schematic diagrams and idealised particle arrangements to illustrate the authors’ understanding of the filter microstructure and its relation to internal stability. For example, Kenney et al. (1985) and Indraratna et al. (2007) estimated constriction sizes using assumed ideal representative particle arrangements. Chang & Zhang (2013) present schematic diagrams illustrating their assumptions of the way fine particles fill the pores between the coarse particles. The prevalence of these hypothetical descriptions in the internal erosion literature has highlighted the need to make direct particle-scale measurements. The present authors know of no prior research that has used high-resolution micro-computed tomography (μCT) data to make three-dimensional particle-scale measurements to quantitatively relate fabric (microstructure) to a well-established internal stability criterion.

Previously Shire & O’Sullivan (2013) used the discrete-element method (DEM) and found a relationship between the internal structure and the Kézdi (1979) criterion for internal stability. The DEM study considered idealised spherical particles; here the research has been extended by using μCT to look at the internal structure of real granular materials.

EXPERIMENTAL METHOD

Quantifying internal stability

As in Shire & O’Sullivan (2013), the internal stability of soils is quantified using the \( \frac{D_{15}}{d_{50}} \) ratio, as proposed by Kézdi (1979). The PSD is divided into coarser and finer fractions. The \( D_{15} \) of the coarse fraction is compared with the \( d_{50} \) of the fine fraction. The ‘finer’ and ‘coarser’ fractions are selected by systematically splitting the PSD at different points along the curve (or at the gap in the case of a gap-graded soil). A soil is considered to be internally stable if \( \frac{D_{15}}{d_{50}} \geq 4 \).

Laboratory sample preparation

Referring to Fig. 1, three specimens of sand with different gradings and different levels of internal stability were created: one well-graded specimen (WG), and two gap-graded specimens (G1 and G2). The gradings were selected to include stable and unstable samples and were achieved by mixing four laboratory sands (quarried sand and processed Leighton Buzzard sand). Table 1 summarises the mixes used and Fig. 1 gives the overall gradings of each sample as measured with standard laboratory sieving and the QicPic laser scanning approach (Witt et al., 2004). For the laser scanning approach three diameter definitions were used. The Feret diameters are the distances between two tangents to a two-dimensional outline of the particle and the maximum and minimum Feret diameters from the laser scanning are denoted \( p_{\text{F}}^{\text{MS}} \) and \( p_{\text{F}}^{\text{MAX}} \).
Table 1. Descriptors for Leighton Buzzard sand and specimen grading constituents by dry mass

<table>
<thead>
<tr>
<th>Supplier grading:</th>
<th>Mesh numbers (BS 410 (BSI, 1986))</th>
<th>% Smaller by volume</th>
<th>% Smaller by volume</th>
<th>% Smaller by volume</th>
<th>% Smaller by volume</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>2360/1180</td>
<td>1180/600</td>
<td>600/300</td>
<td>300/150</td>
<td></td>
</tr>
<tr>
<td>Well graded</td>
<td>7/14</td>
<td>14/25</td>
<td>25/52</td>
<td>52/100</td>
<td></td>
</tr>
<tr>
<td></td>
<td>20%</td>
<td>48%</td>
<td>32%</td>
<td>0%</td>
<td></td>
</tr>
<tr>
<td>Gap-graded 1</td>
<td>86%</td>
<td>0%</td>
<td>2%</td>
<td>12%</td>
<td></td>
</tr>
<tr>
<td>Gap-graded 2</td>
<td>73%</td>
<td>2%</td>
<td>1%</td>
<td>24%</td>
<td></td>
</tr>
</tbody>
</table>

Micro-computed tomography scans

Micro-computed tomography uses the different absorption of X-rays by the materials in a sample to generate a

Resin impregnation

In \( \mu \text{CT} \) the resolution of the images obtained is inversely proportional to the size of the scanned sample (e.g. Fonseca, 2011). To enable analysis of the inter-particle contacts a resolution of 10 \( \mu \text{m} \) was needed and so the sample size for scanning was restricted to 9 mm diameter. To allow these 9 mm cores to be extracted without disturbing the fabric, the specimens were impregnated with a low-viscosity epoxy resin. The resin used was Epo-Tek 301, which has been successfully used in other studies to examine the fabric of cohesionless specimens (e.g. Palmer & Barton, 1986; Jang et al., 1999; Fonseca, 2011). This resin has a relatively low viscosity of 100 mPa s at 25°C, compared to a typical epoxy resin viscosity of 300 mPa s (Jang et al., 1999). Another important characteristic of this resin is the fact that it cures at room temperature, which avoids differential thermal stresses and volume changes related to variations in temperature that generally occur with thermo-curing resins.

The impregnation was carried out under a suction of 50 kPa applied at the top of the sample. The resin was drawn into the base of the specimen from two feed lines from the resin reservoir under the combined influence of hydrostatic head and a suction applied through the top drainage line, while maintaining the stress state at \( p' = 300 \) kPa and \( q = 276 \) kPa, as depicted in Figs 2(b) and 2(c). Once fully impregnated, the suction was removed and the specimen was left to cure under the maintained stress state for 24 h before being extracted from the cell. A central core along the full specimen, 9 mm in diameter, was drilled from each sample. The internal topology of the three specimens was inspected using three-dimensional images from \( \mu \text{CT} \) scans taken at three different locations, namely, the middle, top and base of the cylindrical core.

Fig. 1. Comparison of particle size distributions obtained using sieve analysis, laser scanning and \( \mu \text{CT} \): (a) sample WG; (b) sample G1; (c) sample G2

respectively. The diameter derived from the equivalent circle (EQPC\( ^{LS} \)) is given by the diameter of a circle that has the same area as the projection area of the particle (Sympatec, 2012).

The samples were dry pluviated into a latex membrane lined mould. The two gap-graded samples were made up in 200 g batches to minimise segregation. Sample segregation was not deemed an issue with the well-graded sample and it was created as a single 2 kg batch. Specimen densification and levelling of the upper surface were performed by gentle vibration of the mould at half-way and full fill points of specimen pluviation.

The samples were tested dry, as in the work of Fonseca et al. (2013a). The stress path followed by the samples is illustrated in Fig. 2(a). They were first taken to an effective isotropic stress of 50 kPa (point B, Fig. 2(a)). Assuming an effective angle of shearing resistance (\( \phi' \)) of 35° for the samples and assuming \( K_0 = 1 - \sin \phi' \) for normally consolidated sands (Jaky, 1944), the stress path for \( K_0 \) compression was determined. Each sample was subject to \( K_0 \) compression until the mean effective stress (\( p' \)) reached 300 kPa (point D, Fig. 2(a)).
three-dimensional image dataset of volumetric pixels (voxels), where each voxel is assigned an intensity level corresponding to the attenuation of X-rays at the corresponding point in the sample. Hasan & Alshibli (2010), Hall et al. (2010) and Fonseca et al. (2013a, 2013b) have all demonstrated the applicability of \(\mu\)CT in soil mechanics. Homberg et al. (2012) considered the potential to apply \(\mu\)CT to examine the void space in internal erosion research. The scans were carried out using a Phoenix nanotom (GE Measuring and Control). This system uses a small focal spot that ensures very little geometric unsharpness and so good image resolution. The suitability of this system to obtain high-quality images of silica sands at a few microns resolution is reported in Fonseca (2011). The images acquired have a voxel size of 10 \(\mu\)m, that is 0.005\(d_{50}\)–0.01\(d_{50}\), considerably higher when compared with key earlier geotechnical studies (see Fonseca et al. (2013b) for details). As indicated in Table 2, the \(d_{10}\) for all samples is represented by not less than 25 voxels.

### IMAGE PROCESSING AND ANALYSIS

#### Image binarisation

Micro-tomography images contain a rich amount of information concerning the internal microstructure of the specimens. Following image reconstruction, a three-dimensional...
median filter of size $3 \times 3 \times 3$ was applied to reduce noise and smooth the image (while preserving the edges). This is particularly relevant for the solid phase that contains a less homogeneous distribution of the X-ray attenuation. Fig. 3 shows the line profile of the intensity levels on a representative grain from the ‘WG middle’ sample. The two solid lines represent the greyscale values prior to and after median filtering. The main points to be noted are: the inner greyscale values have been smoothed and the location of the grain’s boundaries, that is the vertical lines, remained unaltered. The histogram of intensity levels in the filtered image had a more distinct bi-modal distribution and lower frequencies for intermediate intensity values, in comparison with the unfiltered data. Thus, the filtering process reduced uncertainties related to the selection of the adequate threshold value.

The threshold value was obtained by analysing the shape of the histogram of the intensity levels of the filtered data. Gaussian curves were fitted to the two peaks of the histogram and the minimum point between them was determined (see Fonseca (2011) for details). This initial threshold value was compared with the values resulting from application of Otsu’s method (Otsu, 1975), the mid value between the two peaks of the histogram and also the threshold value that matches the experimental void ratio (e.g. Chevalier & Otani, 2011). Fig. 3 shows the binary profile, represented by dashed lines, obtained for each method. The initial threshold value was $T = 2353$, Otsu’s value gave $T = 2330$, mid peak value gave $T = 2200$ and the experimental void ratio was $T = 2404$. Despite the small variations in the appropriate threshold value that was indicated from the different methods, whichever method was applied, the edges of the grain remain unchanged. Similar observations were noted for the gap-graded materials; however, the technique that uses the overall void ratio was found unsuitable because of the heterogeneity of the specimens. The output of the thresholding process is a binary image where the pixels were classified as either void or solid phase: that is, the voids were shaded black and particles were shaded white.

Horizontal and vertical cross-sections through the centre of each of the nine samples that were scanned using μCT are presented in Fig. 4; it is clear that all nine μCT specimens include a degree of heterogeneity and that this heterogeneity is more pronounced for the gap-graded material (G1 and G2). It seems that there is segregation in sample G1; there are clearly more fine particles in the base sample of the G1 material and it is easy to imagine that the smaller particles fell through the gaps that appear to exist between the larger grains. Kenney & Lau (1985) propose that in an internally unstable soil there is a primary fabric of stress-transmitting coarse particles and there are loose finer particles in the voids between the primary fabric, which do not carry effective stress and can be moved by seepage. While the amount of information that can be obtained from such two-dimensional sectional images is limited, it does seem that, in the gap-graded material in particular, these criteria are met.
Identification of individual grains
The separation of the interconnected solid phase into individual grains was carried out using a marker-based watershed algorithm (Avizo, VSG). The procedure consisted of: (a) calculating the Euclidian distance map (EDM) on the solid phase of the binary image, (b) identifying the local maxima of the EDM and (c) merging the local maxima to a pre-defined watershed depth value which originates the markers. The number of markers dictates the final number of individual particles coming out of the process, therefore the selected depth value should associate with each individual grain no more than one marker. Small grains, that is grains with lower EDM maxima, require the use of a lower depth value to avoid other local maxima being included in the principal maximum, which would originate under-segmentation of the grains. On the other hand, using lower depth values could cause over-segmentation of very large single grains as it would assign them more than one marker. It was found that a depth value of 6 produced good results for the well-graded samples, whereas for the gap-graded specimens (G1 and G2), the large amount of small grains required a value of 1. The segmentation results for the gap-graded samples were generally good, with the exception of some large grains that, given their high EDM maxima, were associated to more than one marker and therefore slipped into multiple grains. Those grains were easily identified and a Matlab code was written to merge them into a single particle. The high level of detail of the images and the relatively simple particle shape — sub-angular to round with high convexity values (approximately 0.9, from Table 3) — helped minimise the errors often associated with the watershed segmentation.

QUANTITATIVE RESULTS

Particle size measurements
The particle sizes were determined by calculating the principal axes lengths using a Matlab code as described in Fonseca et al. (2011). The procedure consisted of: (a) calculating the Euclidian distance map (EDM) on the solid phase of the binary image, (b) identifying the local maxima of the EDM and (c) merging the local maxima to a pre-defined watershed depth value which originates the markers. The number of markers dictates the final number of individual particles coming out of the process, therefore the selected depth value should associate with each individual grain no more than one marker. Small grains, that is grains with lower EDM maxima, require the use of a lower depth value to avoid other local maxima being included in the principal maximum, which would originate under-segmentation of the grains. On the other hand, using lower depth values could cause over-segmentation of very large single grains as it would assign them more than one marker. It was found that a depth value of 6 produced good results for the well-graded samples, whereas for the gap-graded specimens (G1 and G2), the large amount of small grains required a value of 1. The segmentation results for the gap-graded samples were generally good, with the exception of some large grains that, given their high EDM maxima, were associated to more than one marker and therefore slipped into multiple grains. Those grains were easily identified and a Matlab code was written to merge them into a single particle. The high level of detail of the images and the relatively simple particle shape — sub-angular to round with high convexity values (approximately 0.9, from Table 3) — helped minimise the errors often associated with the watershed segmentation.

<table>
<thead>
<tr>
<th>Sample</th>
<th>Aspect ratio</th>
<th>Convexity</th>
<th>Sphericity</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Feret min.</td>
<td></td>
<td>Sympatec</td>
</tr>
<tr>
<td></td>
<td>Feret max.</td>
<td>Elongation</td>
<td></td>
</tr>
<tr>
<td>Well graded</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>µCT Top</td>
<td>0.792</td>
<td>0.762</td>
<td>0.869</td>
</tr>
<tr>
<td>Middle</td>
<td>0.793</td>
<td>0.766</td>
<td>0.881</td>
</tr>
<tr>
<td>Base</td>
<td>0.790</td>
<td>0.765</td>
<td>0.881</td>
</tr>
<tr>
<td>QicPic</td>
<td>0.758</td>
<td></td>
<td>0.960</td>
</tr>
<tr>
<td>Gap-graded 1</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>µCT Top</td>
<td>0.791</td>
<td>0.745</td>
<td>0.888</td>
</tr>
<tr>
<td>Middle</td>
<td>0.808</td>
<td>0.766</td>
<td>0.869</td>
</tr>
<tr>
<td>Base</td>
<td>0.791</td>
<td>0.771</td>
<td>0.875</td>
</tr>
<tr>
<td>QicPic</td>
<td>0.749</td>
<td></td>
<td>0.970</td>
</tr>
<tr>
<td>Gap-graded 2</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>µCT Top</td>
<td>0.797</td>
<td>0.793</td>
<td>0.847</td>
</tr>
<tr>
<td>Middle</td>
<td>0.797</td>
<td>0.758</td>
<td>0.876</td>
</tr>
<tr>
<td>Base</td>
<td>0.788</td>
<td>0.771</td>
<td>0.874</td>
</tr>
<tr>
<td>QicPic</td>
<td>0.752</td>
<td></td>
<td>0.967</td>
</tr>
</tbody>
</table>

Analysis of the contacts between particles
Shire & O’Sullivan (2013) proposed that there is a relationship between the average number of contacts per particle (i.e. the coordination number) and the stability of the material as assessed using the Kézdi criterion. For each of the nine µCT samples the Kézdi ratio, \( \frac{N_c}{N_p} \), was calculated using the intermediate axis length \( b \). The µCT data were analysed using the approaches documented in Fonseca (2011). The coordination number, \( Z \), is defined as \( Z = 2N_c/N_p \), where \( N_c \) and \( N_p \) are the numbers of contacts and particles respectively and each contact is shared between two particles. For each µCT image, the \( Z \) values were calculated using all particles and all contacts detected in the image. Fig. 5 is a plot of \( Z \) against \( \frac{D_{15}/d_{85}}{\max} \) for the µCT data and the DEM dataset presented in Shire & O’Sullivan (2013). Considering the µCT data, the general trend observed by Shire & O’Sullivan is captured; however, the coordination number values are significantly higher. There is slightly more scatter for the sand.

The higher coordination numbers observed in the experimental study are a consequence of the non-spherical particles present in the real sand. Irregular shaped grains are prone to form a higher number of contacts, moreover, these particles can have more than one particle–particle contact. Sample G2 is at the cusp of being overfilled. In an overfilled
between the coarse particles. In fact, the partial coordination
whether the fine particles completely fill the voids
particularly the size ratio between coarse and fine particles
no clear trend. This suggests that in addition to void ratio,
crease in void ratio. The gap-graded samples, however, show
well-graded soil an increase in
set al. 2013b), indicating that for a
potential of using μCT data to verify conclusions developed using DEM results. The work also shows that the
correlation between void ratio and coordination number is
highly dependent on the shape of the PSD curve. Pre-
existing datasets all considered well-graded or uniformly
graded materials and found exponential relationships be-
between void ratio and Z. Here the gap-graded materials
clearly do not follow that trend.

CONCLUSION
This study has used μCT to experimentally examine con-
cclusions of a DEM-based study by Shire & O’Sullivan (2013). A key finding of this earlier study was that the
coordination number of a filter reduces monotonically as the
stability of the filter decreases when the stability is quanti-
fied according to Kézdi (1979). The experimental data con-
firm this finding and show that for real materials the
magnitude of the coordination number is considerably higher
when compared with systems of spheres. This study rein-
forces the potential of using μCT data to verify conclusions
developed using DEM results. The work also shows that the
correlation between void ratio and coordination number is
highly dependent on the shape of the PSD curve. Pre-
existing datasets all considered well-graded or uniformly
graded materials and found exponential relationships be-
between void ratio and Z. Here the gap-graded materials
clearly do not follow that trend.

ACKNOWLEDGEMENTS
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studies of Mr Thomas Shire are funded by the Engineering
and Physical Sciences Research Council (doctoral training
grant).

NOTATION

\( a \)  major principal axis length
\( b \)  intermediate principal axis length
\( C_u \)  coefficient of uniformity (\( = d_{90}/d_{10} \))
\( C_i \)  coefficient of curvature (\( = d_{60}^2/(d_{85} \times d_{15}) \))
\( c \)  minor principal axis length
\( D_{55} \)  coarse fraction of the PSD curve
\( d_{10} \)  particle size for which 10% of particles are smaller
\( d_{30} \)  particle size for which 30% of particles are smaller
\( d_{60} \)  median particle size
\( d_{85} \)  particle size for which 85% of particles are smaller
\( d_{95} \)  fine fraction of the PSD curve
\( EQPC_{LS} \)  diameter of the circle of equal projection area from laser
scanning data
\( P_{1S}^{MN} \)  minimum Feret diameter from laser scanning data
\( P_{1S}^{MAX} \)  maximum Feret diameter from laser scanning data
\( K_o \)  coefficient of earth pressure at rest
\( N_c \)  number of contacts
\( N_p \)  number of particles
\( \rho' \)  mean normal effective stress (\( = (\sigma_1 + 2\sigma_3)/3 \))
\( q \)  deviatoric stress (\( = \sigma_1 - \sigma_3 \))
\( Z \)  coordination number
\( \sigma_1' \)  axial effective stress
\( \sigma_3' \)  radial effective stress
\( \phi' \)  effective angle of shearing resistance

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