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TECHNICAL NOTE

Microstructural analysis of sands with varying degrees of internal stability

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Internal erosion involves the migration of particles through a geotechnical structure. Internal erosion poses a significant hazard to embankment dams and flood embankments. The fundamental mechanisms operate at the particle scale and a thorough understanding of these mechanisms can inform the filter design and specification process and reduce the hazard that internal erosion is known to pose to many engineered embankment structures. Engineers have long acknowledged the importance of the grain scale interactions, but until recently, explanations of the mechanisms have been purely hypothetical, as direct observation of the internal structure of filters was not possible. Recent research has used the discrete-element method to establish a particle-scale basis for Kézdi's filter internal stability criterion. The discrete-element method can provide significant useful data on soil microstructure, so a discrete-element method model is inherently ideal. This study therefore examines a number of real sand samples with varying degrees of internal stability at the particle scale using high-resolution micro-computed tomography. The correlation between coordination number and internal stability is con-firmed, with the coordination number values being significantly higher for the real material.

KEYWORDS: erosion; filters; laboratory tests

INTRODUCTION

Embankment dam filters prevent erosion of a low-permeability clay core due to water seepage. Filters should be erosion resistant or 'internally stable'. Internally unstable filters are susceptible to suffusion, an internal erosion mechanism that occurs when the finer particles within the filter are carried downstream in the direction of seepage flow (e.g. Fell & Fry, 2007). Three factors contribute to internal erosion: (*a*) material susceptibility, (*b*) hydraulic susceptibility and (*c*) mechanical susceptibility (Brown & Bridle, 2009). Design guidelines focus on the material susceptibility and current guidelines for assessing the internal stability of filters are mainly derived empirically and consider the shape of the particle size distribution (PSD) curve (e.g. Kézdi, 1979; Kenney & Lau, 1985, 1986; Burenkova, 1993).

The initiation and propagation of internal erosion is a particle-scale phenomenon. Prior discussions on internal erosion have often included schematic diagrams and idealised particle arrangements to illustrate the authors' understanding of the filter microstructure and its relation to internal stability. For example, Kenney *et al.* (1985) and Indraratna *et al.* (2007) estimated constriction sizes using assumed ideal representative particle arrangements. Chang & Zhang (2013) present schematic diagrams illustrating their assumptions of the way fine particles fill the pores between the coarse particles. The prevalence of these hypothetical descriptions in the internal erosion literature has highlighted the need to make direct particle-scale measurements. The present authors know of no prior research that has used high-resolution micro-computed tomography (μ CT) data to make three-

dimensional particle-scale measurements to quantitatively relate fabric (microstructure) to a well-established internal stability criterion.

Previously Shire & O'Sullivan (2013) used the discreteelement method (DEM) and found a relationship between the internal structure and the Kézdi (1979) criterion for internal stability. The DEM study considered idealised spherical particles; here the research has been extended by using μ CT to look at the internal structure of real granular materials.

EXPERIMENTAL METHOD

Quantifying internal stability

As in Shire & O'Sullivan (2013), the internal stability of soils is quantified using the $(D'_{15}/d'_{85})_{\text{max}}$ ratio, as proposed by Kézdi (1979). The PSD is divided into coarser and finer fractions. The D'_{15} of the coarse fraction is compared with the d'_{85} of the fine fraction. The 'finer' and 'coarser' fractions are selected by systematically splitting the PSD at different points along the curve (or at the gap in the case of a gap-graded soil). A soil is considered to be internally stable if $(D'_{15}/d'_{85})_{\text{max}} < 4$.

Laboratory sample preparation

Referring to Fig. 1, three specimens of sand with different gradings and different levels of internal stability were created; one well-graded specimen (WG), and two gap-graded specimens (G1 and G2). The gradings were selected to include stable and unstable samples and were achieved by mixing four laboratory sands (quarried sand and processed Leighton Buzzard sand). Table 1 summarises the mixes used and Fig. 1 gives the overall gradings of each sample as measured with standard laboratory sieving and the QicPic laser scanning approach (Witt *et al.*, 2004). For the laser scanning approach three diameter definitions were used. The Feret diameters are the distances between two tangents to a two-dimensional outline of the particle and the maximum and minimum Feret diameters from the laser scanning are denoted F_{MAX}^{LS} and F_{MAX}^{LS}

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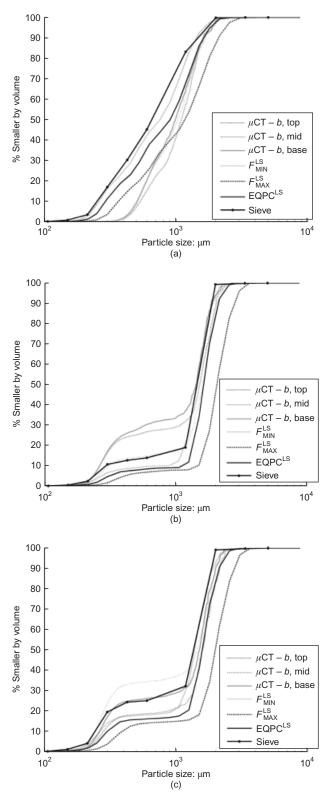


Fig. 1. Comparison of particle size distributions obtained using sieve analysis, laser scanning and μ CT: (a) sample WG; (b) sample G1; (c) sample G2

respectively. The diameter derived from the equivalent circle (EQPC^{LS}) is given by the diameter of a circle that has the same area as the projection area of the particle (Sympatec, 2012).

The samples were dry pluviated into a latex membrane lined mould. The two gap-graded samples were made up in 200 g batches to minimise segregation. Sample segregation was not deemed an issue with the well-graded sample and it was created as a single 2 kg batch. Specimen densification and levelling of the upper surface were performed by gentle vibration of the mould at half-way and full fill points of specimen pluviation.

The samples were tested dry, as in the work of Fonseca *et al.* (2013a). The stress path followed by the samples is illustrated in Fig. 2(a). They were first taken to an effective isotropic stress of 50 kPa (point B, Fig. 2(a)). Assuming an effective angle of shearing resistance (ϕ') of 35° for the samples and assuming $K_0 = 1 - \sin\phi'$ for normally consolidated sands (Jaky, 1944), the stress path for K_0 compression was determined. Each sample was subject to K_0 compression until the mean effective stress (p') reached 300 kPa (point D, Fig. 2(a)).

Resin impregnation

In μ CT the resolution of the images obtained is inversely proportional to the size of the scanned sample (e.g. Fonseca, 2011). To enable analysis of the inter-particle contacts a resolution of 10 µm was needed and so the sample size for scanning was restricted to 9 mm diameter. To allow these 9 mm cores to be extracted without disturbing the fabric, the specimens were impregnated with a low-viscosity epoxy resin. The resin used was Epo-Tek 301, which has been successfully used in other studies to examine the fabric of cohesionless specimens (e.g. Palmer & Barton, 1986; Jang et al., 1999; Fonseca, 2011). This resin has a relatively low viscosity of 100 mPa s at 25°C, compared to a typical epoxy resin viscosity of 300 mPas (Jang et al., 1999). Another important characteristic of this resin is the fact that it cures at room temperature, which avoids differential thermal stresses and volume changes related to variations in temperature that generally occur with thermo-curing resins.

The impregnation was carried out under a suction of 50 kPa applied at the top of the sample. The resin was drawn into the base of the specimen from two feed lines from the resin reservoir under the combined influence of hydrostatic head and a suction applied through the top drainage line, while maintaining the stress state at p' = 300 kPa and q = 276 kPa, as depicted in Figs 2(b) and 2(c). Once fully impregnated, the suction was removed and the specimen was left to cure under the maintained stress state for 24 h before being extracted from the cell. A central core along the full specimen, 9 mm in diameter, was drilled from each sample. The internal topology of the three specimens was inspected using three-dimensional images from μ CT scans taken at three different locations, namely, the middle, top and base of the cylindrical core.

Micro-computed tomography scans

Micro-computed tomography uses the different absorption of X-rays by the materials in a sample to generate a

Table 1. Descriptors for Leighton Buzzard sand and specimen grading constituents by dry mass

Mesh numbers (BS 410 (BSI, 1986))	7/14	14/25	25/52	52/100
Supplier grading: μm	2360/1180	1180/600	600/300	300/150
Well graded	20%	48%	32%	0%
Gap-graded 1	86%	0%	2%	12%
Gap-graded 2	73%	2%	1%	24%

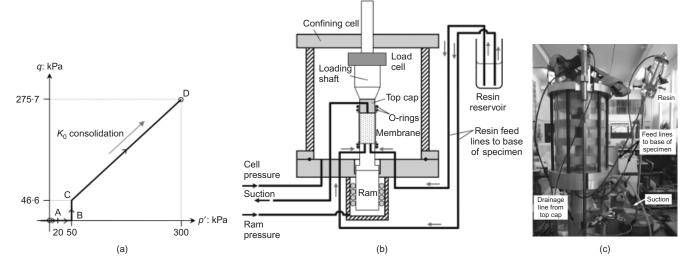


Fig. 2. Laboratory preparation of samples: (a) stress path followed; (b) schematic diagram of impregnation; (c) photograph of resin impregnation in process

three-dimensional image dataset of volumetric pixels (voxels), where each voxel is assigned an intensity level corresponding to the attenuation of X-rays at the corresponding point in the sample. Hasan & Alshibli (2010), Hall *et al.* (2010) and Fonseca *et al.* (2013a, 2013b) have all demonstrated the applicability of μ CT in soil mechanics. Homberg *et al.* (2012) considered the potential to apply μ CT to examine the void space in internal erosion research.

The scans were carried out using a Phoenix nanotom (GE Measuring and Control). This system uses a small focal spot that ensures very little geometric unsharpness and so good image resolution. The suitability of this system to obtain high-quality images of silica sands at a few microns resolu-

tion is reported in Fonseca (2011). The images acquired have a voxel size of 10 μ m, that is $0.005d_{50}-0.01d_{50}$, considerably higher when compared with key earlier geotechnical studies (see Fonseca *et al.* (2013b) for details). As indicated in Table 2, the d_{10} for all samples is represented by not less than 25 voxels.

IMAGE PROCESSING AND ANALYSIS *Image binarisation*

Micro-tomography images contain a rich amount of information concerning the internal microstructure of the specimens. Following image reconstruction, a three-dimensional

		<i>d</i> ₁₀ : μm	d ₅₀ : μm	d ₉₀ : μm	C_{u}	C_z
ell graded						
ave(<i>a</i> , <i>b</i> , <i>c</i>)	Тор	518.0	1040.4	1697.6	2.28	0.89
	Middle	520.6	1009.2	1444.9	2.16	0.96
	Base	521.7	999.7	1582-2	2.18	0.86
QicPic (Fere	t min.)	260.6	756.5	1432.7	3.53	0.89
QicPic (Fere		366.0	1168.6	2106.7	3.75	0.95
Sieve		254.8	673.8	1516.8	3.24	0.85
ap-graded 1						
ave(a,b,c)	Тор	312.4	1592.5	2048.1	5.41	3.81
	Middle	268.5	1560.1	1993.7	6.18	2.91
	Base	268.4	1449.4	1871.7	5.75	0.88
QicPic (Feret min.)		825.0	1436.6	1921.6	1.84	1.32
QicPic (Feret max.)		1472.3	2059.8	2819.5	1.48	1.04
Sieve		293.2	1496.8	1904.7	5.45	3.57
p-graded 2						
ave(a,b,c)	Тор	250.5	1471.4	2075-4	6.41	0.31
	Middle	249.5	1362.5	1734.5	5.85	0.31
	Base	251.9	1405.9	1778.9	6.00	0.30
QicPic (Fere	t min.)	265.6	1385.5	1909.1	5.54	3.59
QicPic (Fere		391.6	1993.7	2790.1	5.41	3.57
Sieve	/	245.6	1397.7	1887.9	6.19	2.67

Table 2. Particle diameters, coefficients of uniformity and curvature from μ CT (ave(a,b,c) is the arithmetic mean of the major principle axis, a, intermediate principle axis, b, and minor principle axis, c), QicPic analysis and manual sieving data

median filter of size $3 \times 3 \times 3$ was applied to reduce noise and smooth the image (while preserving the edges). This is particularly relevant for the solid phase that contains a less homogeneous distribution of the X-ray attenuation. Fig. 3 shows the line profile of the intensity levels on a representative grain from the 'WG middle' sample. The two solid lines represent the greyscale values prior to and after median filtering. The main points to be noted are: the inner greyscale values have been smoothed and the location of the grain's boundaries, that is the vertical lines, remained unaltered. The histogram of intensity levels in the filtered image had a more distinct bi-modal distribution and lower frequencies for intermediate intensity values, in comparison with the unfiltered data. Thus, the filtering process reduced uncertainties related to the selection of the adequate threshold value.

The threshold value was obtained by analysing the shape of the histogram of the intensity levels of the filtered data. Gaussian curves were fitted to the two peaks of the histo-

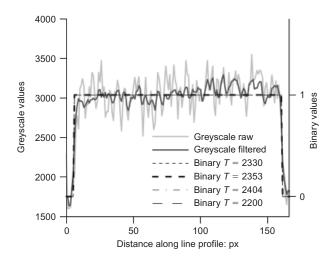


Fig. 3. Line profile along a representative grain

gram and the minimum point between them was determined (see Fonseca (2011) for details). This initial threshold value was compared with the values resulting from application of Otsu's method (Otsu, 1975), the mid value between the two peaks of the histogram and also the threshold value that matches the experimental void ratio (e.g. Chevalier & Otani, 2011). Fig. 3 shows the binary profile, represented by dashed lines, obtained for each method. The initial threshold value was T = 2353, Otsu's value gave T = 2330, mid peak value gave T = 2200 and the experimental void ratio was T = 2404. Despite the small variations in the appropriate threshold value that was indicated from the different methods, whichever method was applied, the edges of the grain remain unchanged. Similar observations were noted for the gap-graded materials; however, the technique that uses the overall void ratio was found unsuitable because of the heterogeneity of the specimens. The output of the thresholding process is a binary image where the pixels were classified as either void or solid phase: that is, the voids were shaded black and particles were shaded white.

Horizontal and vertical cross-sections through the centre of each of the nine samples that were scanned using μ CT are presented in Fig. 4; it is clear that all nine μ CT specimens include a degree of heterogeneity and that this heterogeneity is more pronounced for the gap-graded material (G1 and G2). It seems that there is segregation in sample G1; there are clearly more fine particles in the base sample of the G1 material and it is easy to imagine that the smaller particles fell through the gaps that appear to exist between the larger grains. Kenney & Lau (1985) propose that in an internally unstable soil there is a primary fabric of stresstransmitting coarse particles and there are loose finer particles in the voids between the primary fabric, which do not carry effective stress and can be moved by seepage. While the amount of information that can be obtained from such two-dimensional sectional images is limited, it does seem that, in the gap-graded material in particular, these criteria are met.

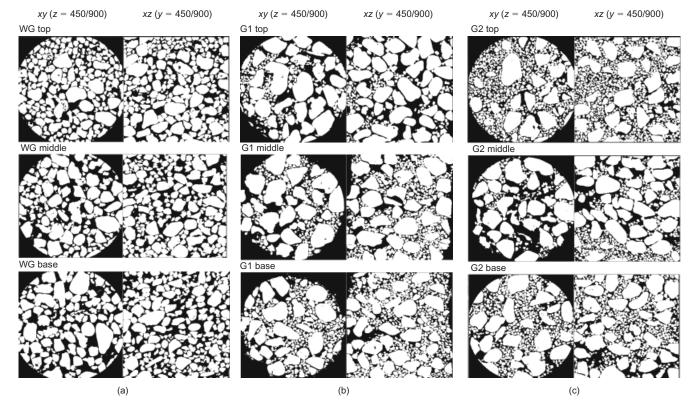


Fig. 4. Sections through μ CT data after image binarisation: (a) WG sample; (b) G1 sample; (c) G2 sample

The separation of the interconnected solid phase into individual grains was carried out using a marker-based watershed algorithm (Avizo, VSG). The procedure consisted of: (a) calculating the Euclidian distance map (EDM) on the solid phase of the binary image, (b) identifying the local maxima of the EDM and (c) merging the local maxima to a pre-defined watershed depth value which originates the markers. The number of markers dictates the final number of individual particles coming out of the process, therefore the selected depth value should associate with each individual grain no more than one marker. Small grains, that is grains with lower EDM maxima, require the use of a lower depth value to avoid other local maxima being included in the principal maximum, which would originate under-segmentation of the grains. On the other hand, using lower depth values could cause oversegmentation of very large single grains as it would assign them more than one marker. It was found that a depth value of 6 produced good results for the well-graded samples, whereas for the gap-graded specimens (G1 and G2), the large amount of small grains required a value of 1. The segmentation results for the gap-graded samples were generally good, with the exception of some large grains that, given their high EDM maxima, were associated to more than one marker and therefore slipped into multiple grains. Those grains were easily identified and a Matlab code was written to merge them into a single particle. The high level of detail of the images and the relatively simple particle shape - sub-angular to round with high convexity values (approximately 0.9, from Table 3) - helped minimise the errors often associated with the watershed segmentation.

QUANTITATIVE RESULTS

Particle size measurements

The particle sizes were determined by calculating the principal axes lengths using a Matlab code as described in Fonseca *et al.* (2012). The segregation effects that are

Table 3. Shape factors from μ CT and QicPic analysis	Table 3.	Shape	factors	from	μCT	and	QicPic	analysis
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apparent from visual observation of Fig. 4 led to a variation in the PSDs obtained for each material. Fig. 1 illustrates the intermediate axes lengths, *b*, for all the μ CT scans. Table 2 gives representative particle diameters for all nine samples. The PSDs differ within each material owing to the segregation effects discussed above. Fonseca *et al.* (2012) found a good match between sieving and the intermediate axis length (*b*), the $F_{\rm MAX}^{\rm LS}$ values were close to the major (maximum) axis length measurement and $F_{\rm MIN}^{\rm LS}$ values slightly exceeded the minor axis length values. Here the various measures of size do not agree as closely as the data presented by Fonseca *et al.* (2012).

Analysis of the contacts between particles

Shire & O'Sullivan (2013) proposed that there is a relationship between the average number of contacts per particle (i.e. the coordination number) and the stability of the material as assessed using the Kézdi criterion. For each of the nine μ CT samples the Kézdi ratio, $(D'_{15}/d'_{85})_{\text{max}}$, was calculated using the intermediate axis length (b). The μ CT data were analysed using the approaches documented in Fonseca (2011). The coordination number, Z, is defined as $Z = 2N_{\rm c}/N_{\rm p}$, where $N_{\rm c}$ and $N_{\rm p}$ are the numbers of contacts and particles respectively and each contact is shared between two particles. For each μ CT image, the Z values were calculated using all particles and all contacts detected in the image. Fig. 5 is a plot of Z against $(D'_{15}/d'_{85})_{\text{max}}$ for the μ CT data and the DEM dataset presented in Shire & O'Sullivan (2013). Considering the μ CT data, the general trend observed by Shire & O'Sullivan is captured; however, the coordination number values are significantly higher. There is slightly more scatter for the sand.

The higher coordination numbers observed in the experimental study are a consequence of the non-spherical particles present in the real sand. Irregular shaped grains are prone to form a higher number of contacts, moreover, these particles can have more than one particle–particle contact. Sample G2 is at the cusp of being overfilled. In an overfilled

		Aspect ratio			Convexity	Sphericity		
		<u>Feret min.</u> Feret max.	Elongation	Flatness		Sympatec (2012)	Sneed & Folk (1958)	Krumbein (1941)
Well graded								
μCT	Top Middle Base		0·792 0·793 0·790	0·762 0·766 0·765	0·869 0·881		0·749 0·761 0·767	0·750 0·763 0·769
QicPic	I	0.758			0.960	0.828		
Gap-graded 1								
μCT	Top Middle Base	-	0·791 0·808 0·791	0.745 0.766 0.771	0.888 0.869 0.875		0.770 0.771 0.749	0.771 0.772 0.744
QicPic		0.749			0.970	0.894		
Gap-graded 2		-						
μCT	Top Middle Base		0·797 0·797 0·788	0.793 0.758 0.771	0·847 0·876 0·874		0.755 0.793 0.755	0·759 0·793 0·759
QicPic	1	0.752			0.967	0.895		

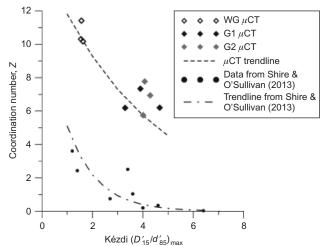


Fig. 5. Comparison of Kézdi criterion with coordination number

sample, the fine particles completely fill the void space between the larger particles and actively transmit stress.

Hasan & Alshibli (2010) and Fonseca et al. (2013b) proposed relationships between the void ratio and the coordination number, with the correlations taking an exponential form, as illustrated in Fig. 6. The variation in void ratio with coordination number for the simulations in Shire & O'Sullivan (2013) is also presented. The WG specimens exhibit values close to what was observed for the well-sorted Reigate sand (Fonseca et al., 2013b), indicating that for a well-graded soil an increase in Z is associated with a decrease in void ratio. The gap-graded samples, however, show no clear trend. This suggests that in addition to void ratio, coordination number is highly dependent on the PSD, in particular the size ratio between coarse and fine particles and whether the fine particles completely fill the voids between the coarse particles. In fact, the partial coordination

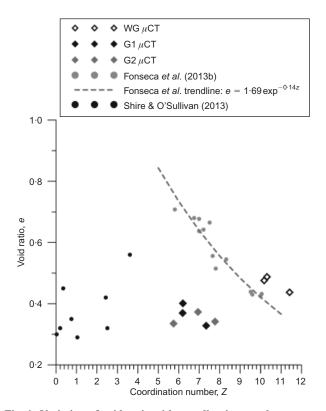


Fig. 6. Variation of void ratio with coordination number

numbers (i.e. large-large, small-small and large-small) such as those used by Minh & Cheng (2012) may be more relevant when considering gap-graded materials. Pinson et al. (1998) found that partial coordination numbers in materials with a bimodal particle size distribution varied dramatically depending on size ratio between small and large and the fines content, and Shire & O'Sullivan (2013) showed that the coarse particles in gap-graded materials play a disproportionate role in force transfer.

CONCLUSION

This study has used μ CT to experimentally examine conclusions of a DEM-based study by Shire & O'Sullivan (2013). A key finding of this earlier study was that the coordination number of a filter reduces monotonically as the stability of the filter decreases when the stability is quantified according to Kézdi (1979). The experimental data confirm this finding and show that for real materials the magnitude of the coordination number is considerably higher when compared with systems of spheres. This study reinforces the potential of using μ CT data to verify conclusions developed using DEM results. The work also shows that the correlation between void ratio and coordination number is highly dependent on the shape of the PSD curve. Preexisting datasets all considered well-graded or uniformly graded materials and found exponential relationships between void ratio and Z. Here the gap-graded materials clearly do not follow that trend.

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NOTATION

- major principal axis length а
- intermediate principal axis length h
- coefficient of uniformity (= d_{60}/d_{10}) $C_{\rm u}$
- coefficient of curvature $(= d_{30}^2/(d_{60} \times d_{10}))$ C_z
- minor principal axis length С
- D_{15} coarse fraction of the PSD curve
- d_{10} particle size for which 10% of particles are smaller
- particle size for which 30% of particles are smaller d_{30}
 - median particle size
- d_{50}
- particle size for which 60% of particles are smaller d_{60}
- fine fraction of the PSD curve d%5
- EQPCLS diameter of the circle of equal projection area from laser scanning data
 - minimum Feret diameter from laser scanning data
 - maximum Feret diameter from laser scanning data MAX
 - coefficient of earth pressure at rest K_0
 - $N_{\rm c}$ number of contacts
 - $N_{\rm p}$ number of particles
 - mean normal effective stress (= $(\sigma'_1 + 2\sigma'_3)/3$) р
 - deviatoric stress (= $\sigma'_1 \sigma'_3$)
 - Ζ coordination number
 - axial effective stress σ'_1
 - σ'_3 radial effective stress
 - effective angle of shearing resistance φ

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